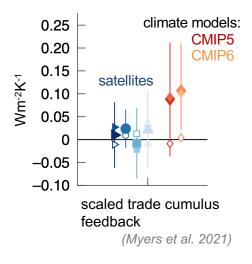
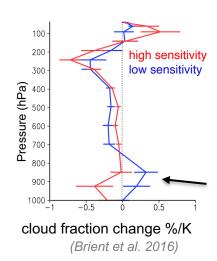
# Strong cloud-circulation coupling explains weak trade cumulus feedback

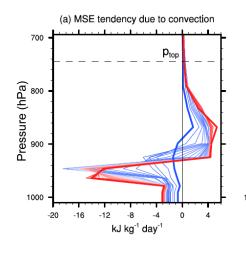
Raphaela Vogel\*, Anna Lea Albright, Jessica Vial, Geet George, Bjorn Stevens, Sandrine Bony

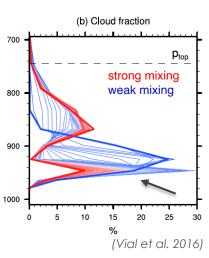


- > the trade-cumulus cloud feedback has remained a major source of uncertainty for climate sensitivity (Bony and Dufresne 2005, Vial et al. 2013, Myers et al. 2021)
- > while many climate models exhibit strong trade cumulus feedbacks, satellite-derived constraints from observed natural variability (Myers et al. 2021, Cesana and del Genio 2021) & large-eddy simulations (Vogel et al. 2016, Radtke et al. 2021) suggest a rather weak feedback
- > In climate models, trade cumulus feedbacks are governed by changes in cloud fraction near cloud base (Vial et al. 2016, Brient et al. 2016)
- > high sensitivity models suggest a desiccation of the lower cloud layer with increasing lower-tropospheric mixing (Vial et al. 2016, Sherwood et al. 2014)



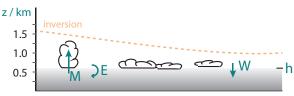




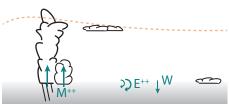


## Mixing-desiccation mechanism – a hypothesis for a strongly positive trade cumulus feedback

#### Base state



**a** Mixing-desiccation mechanism ( $\beta$ <0)



h: sub-cloud layer top

M: mass flux

E: entrainment rate

W: mesoscale vertical velocity

C: cloud-base cloud fraction

R: mean relative humidity

\_\_\_\_\_

- enhanced moisture transport by convection compensated by downward mixing of drier air & evaporation of clouds near cloud base.
- $\rightarrow$  C  $\propto$  R  $\propto$  M<sup> $\beta$ </sup>, with  $\beta$  < 0
- consistent with high-sensitivity climate models & idealized large-eddy simulations of non-precipitating trade cumuli (Sherwood et al. 2014, Rieck et al. 2012)

#### but.....

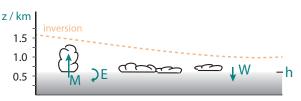
- $M_{act} = Cac_t wac_t$ , mostly governed by area fraction of active clouds  $C_{act}$  (~50% of total C)  $\Rightarrow \beta > 0$
- substantial variability in W observed in the trades (Bony & Stevens 2019, George et al. 2021)
- never tested with observations



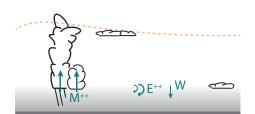
### Mixing-desiccation mechanism – a hypothesis for a strongly positive trade cumulus feedback







**a** Mixing-desiccation mechanism ( $\beta$ <0)



h: sub-cloud layer top

M: mass flux

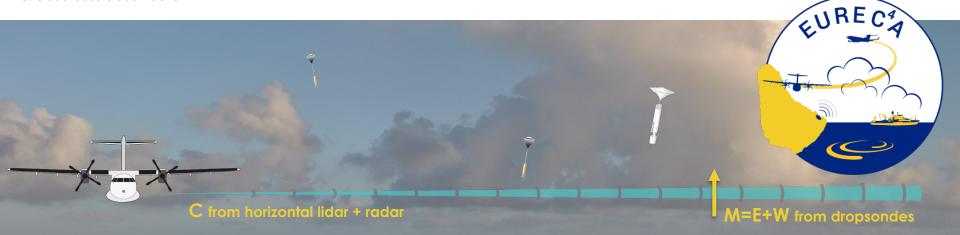
E: entrainment rate

W: mesoscale vertical velocity

C: cloud-base cloud fraction

 $\frac{Dh}{Dt} = E + W - M$ 

- enhanced moisture transport by convection compensated by downward mixing of drier air & evaporation of clouds near cloud base.
- $\rightarrow$  C  $\propto$  R  $\propto$  M<sup> $\beta$ </sup>, with  $\beta$  < 0
- consistent with high-sensitivity climate models & idealized large-eddy simulations of non-precipitating trade cumuli (Sherwood et al. 2014, Rieck et al. 2012)

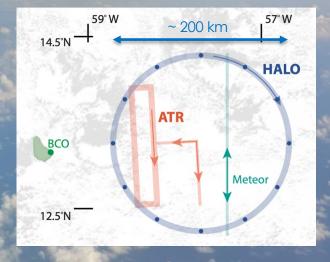


### EUREC<sup>4</sup>A field campaign

(Bony et al. 2017, Stevens et al. 2021)

- Jan-Feb 2020
- 4 aircraft & ships, drones, BCO...
- goal: test mixing-desiccation hypothesis
- Clouds @Barbados representative for entire trade-wind belt (Medeiros & Nuijens 2016)





mass flux estimation from dropsonde measurements

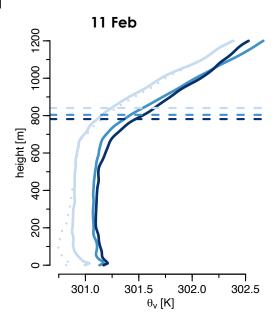
# Mass flux estimation using EUREC<sup>4</sup>A dropsondes

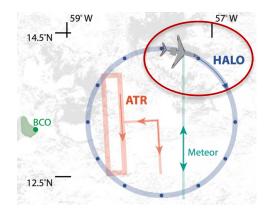
$$M = E + W - \frac{\partial h}{\partial t} \sqrt{h \cdot \nabla h}$$

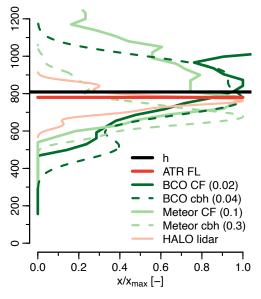
~ Mact=aact Wact (Vogel et al. 2020)

> sub-cloud layer top h

- target: max. cloud-base cloud fraction level
- definition:  $\theta_v(h) \ge \overline{\theta_v} + \epsilon$ , with  $\epsilon = 0.2K$







# Mass flux estimation using EUREC<sup>4</sup>A dropsondes

$$M = E + W - \frac{\partial h}{\partial t} = \nabla_h \cdot \nabla h$$

> sub-cloud layer top h

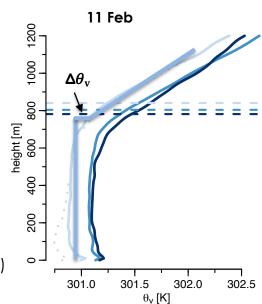
- target: max. cloud-base cloud fraction level
- definition:  $\theta_v(h) \ge \overline{\theta_v} + \epsilon$ , with  $\epsilon = 0.2K$

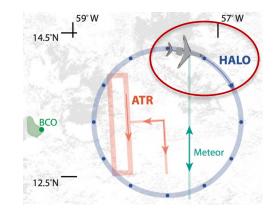
> entrainment rate E:

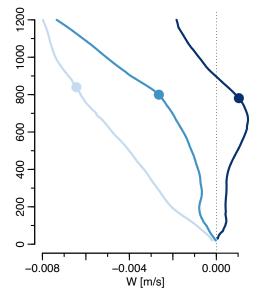
$$E=rac{A_{
m e}\,\overline{w' heta'_{
m v}}|_{
m s}}{\Delta heta_{
m v}}$$
 , with  $A_{
m e}$  = 0.43 (Albright et al., 2022)

> mesoscale vertical velocity W at h: from regression method (Bony & Stevens 2019)

>> target scale: 3-circle averages (~3h, 200 km)

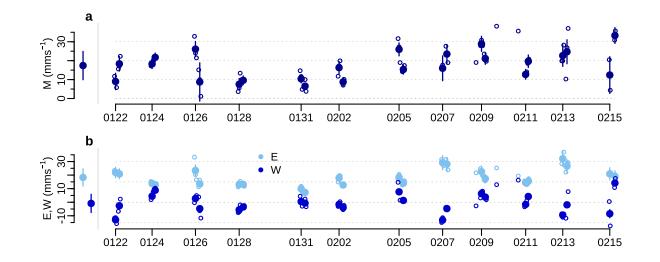






# First observations of convective mixing at the mesoscale

- M and E robust to changes in estimation procedure and consistent with independent data
- on average, M~E
- but on shorter timescales, E & W contribute almost equally to variability in M



# Cloud-base cloud fraction

horizontally-staring 355nm ALIAS lidar

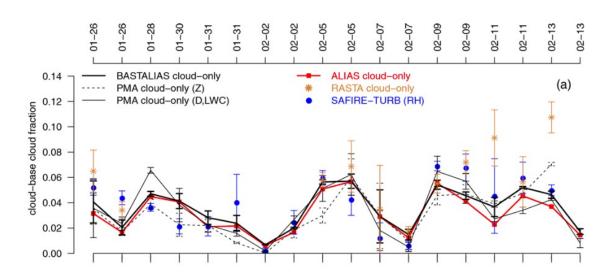


Chazette et al. 2020)

horizontally-staring 94GHz BASTA Doppler radar



(Delanoëet al. 2016)

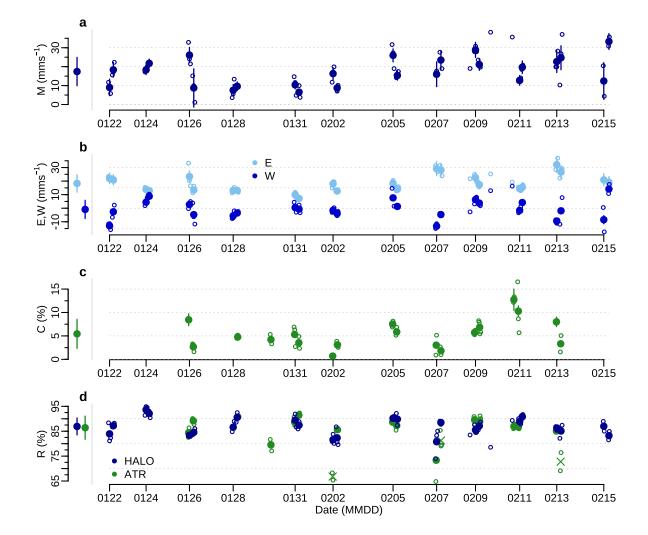


very good agreement among different instruments (Bony et al. 2022)

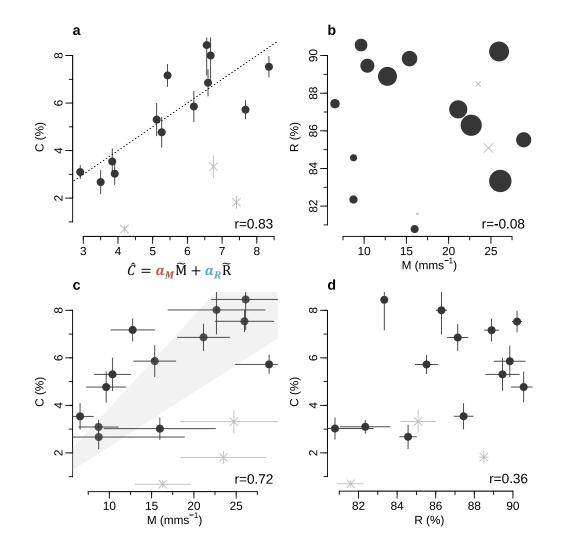


# First observations of M, C and RH co-variations

- C is both small and highly variable
- R is robustly around 86%
- 3 circle-sets with inconsistent sampling neglected



Do we find evidence for the mixing-desiccation mechanism in the EUREC<sup>4</sup>A data?



M: mass flux
E: entrainment rate
W: mesoscale vertical velocity
C: cloud-base cloud fraction
R: mean relative humidity

W & E contribute equally to variability in M, but have opposing relations to R

→ negligible desiccation effect of M!

M alone explains 50% of C variability

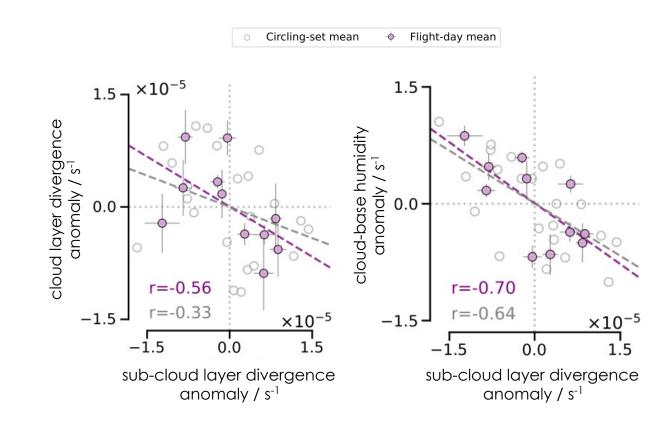
dynamical control through M overwhelms thermodynamic control through R  $\rightarrow \alpha_M/\alpha_R \sim 1.8$ 

EUREC<sup>4</sup>A data refute mixing-desiccation mechanism

# Ubiquity of SMOCS\* and their influence on moisture variance in the trades (George et al. 2022, in review)

\*Shallow Mesoscale Overturning Circulations

- anti-correlation between divergence in the sub-cloud and cloud layers
- Sub-cloud convergence correlated with moister subcloud and cloud-base layers
- ERA5: SMOCs are elongated features of ~100-200 km and cover ~58% of domain



- 4 CMIP5 and 6 CMIP6 models (Taylor et al. 2012, Eyring et al. 2016)

- AMIP 1979-2008 & AMIP+4K (uniform warming)

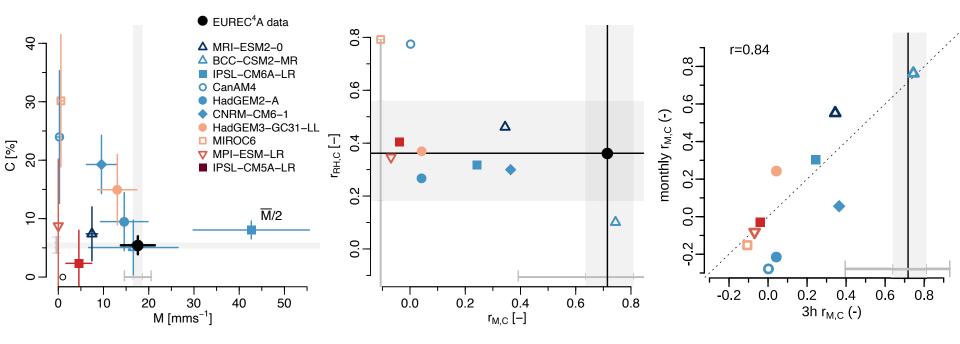
- Winter months (DJFM)

- subhourly output at selected sites from CFMIP (Webb et al. 2017): BCO, BOMEX, EUREC<sup>4</sup>A, NTAS

- monthly outputs over 60W-44W, 11N-16N

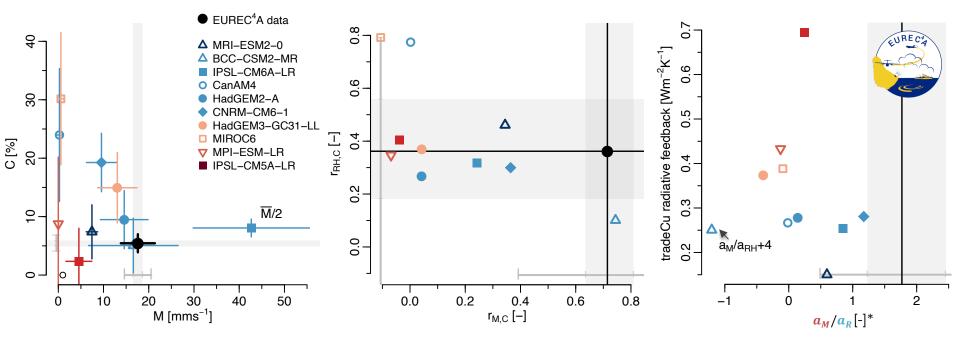
How consistent is the present generation of climate models with our observations?

# Models underestimate strong cloud-circulation coupling



Magnitude, variability, and coupling of M, C and R in CFMIP models differs drastically from EUREC<sup>4</sup>A data Underlying fast physical processes that couple M, R and C in the models are largely time-scale invariant

Process-based constraints render strongly positive trade cumulus feedbacks implausible



Magnitude, variability, and coupling of M, C and R in CFMIP models differs drastically from EUREC<sup>4</sup>A data

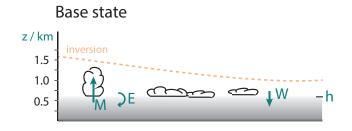
Underlying fast physical processes that couple M, R and C in the models are largely time-scale invariant

Models with largest positive feedback represent refuted mixing-desiccation mechanism and particularly exaggerate variability of C and coupling of C to R instead of M (small  $a_M/a_R$ )

\* $a_M/a_R$  from  $\hat{C} = a_M \tilde{M} + a_R \tilde{R}$ 

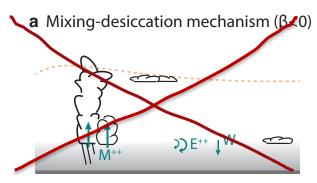
conclusions

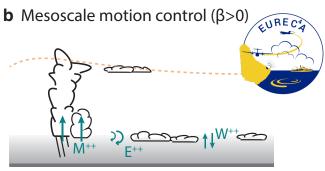
### **Conclusions**



EUREC<sup>4</sup>A emphasizes dynamic factors—convective and mesoscale motions—as dominant controls of cloudiness, rather than thermodynamic factors related to the mixing-desiccation mechanism.

By refuting the mixing-desiccation mechanism, the EUREC<sup>4</sup>A data...





- ... refute an important mechanism for a strongly positive trade cumulus feedback and thus a critical line of evidence for a large climate sensitivity (Stevens et al. 2016)
- ... render climate models with strong positive feedbacks implausible
- ... both support (Myers et al. 2021, Vogel et al. 2016) and explain at the process scale a weak trade cumulus feedback

paper accepted in Nature, preprint: <a href="https://doi.org/10.1002/essoar.10512547.1">https://doi.org/10.1002/essoar.10512547.1</a>